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## A TERRAIN-BASED SITE CONDITIONS MAP OF CALIFORNIA WITH IMPLICATIONS FOR THE CONTIGUOUS UNITED STATES

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## ABSTRACT

We present an approach based on geomorphometry to predict material properties and characterize site conditions using the  $V_{S30}$  (timeaveraged shear-wave velocity to a depth of 30 meters) parameter. Our framework consists of an automated terrain classification scheme based on taxonomic criteria (slope gradient, local convexity and surface texture) that systematically identifies 16 terrain types from 1-km spatial resolution (30 arc sec) Shuttle Radar Topography Mission digital elevation models (SRTM DEMs). Using 853  $V_{S30}$ values from California, we apply a simulation-based statistical method to determine the mean  $V_{S30}$  for each terrain type in California. We then compare the  $V_{S30}$  values to models based on individual proxies, such as mapped surface geology and topographic slope, and show that our systematic terrain-based approach consistently performs better than semi-empirical estimates based on individual proxies. To further evaluate our model, we apply our California-based estimates to terrains of the contiguous U.S. Comparisons of our estimates to 325  $V_{S30}$  measurements outside of California, as well as estimates based on the topographic slope model, indicate our method to be statistically robust and more accurate. Our approach, thus, provides an objective and robust method for extending estimates of  $V_{S30}$  where in-situ measurements are sparse or not readily available.

## INTRODUCTION

To assess site conditions, key factors including layer thickness and impedance (the product of the density of the material and the velocity of the propagating wave) are considered. Because shear-wave velocity ( $V_S$ ) correlates with soil rigidity and is characterized by higher variability than density,  $V_S$  has generally been accepted as an appropriate measure of soil conditions (e.g., Borcherdt, 1970; Fumal, 1978; Aki, 1988; Borcherdt and Glassmoyer, 1992; Boore, 2006). Although variations in material properties at depths of tens to hundreds of meters below the Earth's surface are known to significantly influence ground motions, deeper variations are also important (e.g., Anderson *et al.*, 1996; Frankel *et al.*, 2002; Boore, 2004; Holzer *et al.*, 2005). Nevertheless, a number of studies (e.g., Borcherdt *et al.*, 1991; Boore *et al.*, 1993; Borcherdt, 1994; Dobry *et al.*, 2000; Boore *et al.*, 1994; 1997; Wills and Silva, 1998) have found good correlation between observed site amplification and  $V_{S30}$ , or the time-averaged shear-wave velocity to a depth of 30 meters below the surface. As a result, the  $V_{S30}$  parameter has since been widely adopted as the key parameter for site characterization.

When direct measurements are prohibitive, sparse or not readily available, the use of proxies, i.e., geologic properties and/or units, topographic slope, terrain types, etc., to infer  $V_S$  or  $V_{S30}$  is typically applied. For example, a number of proxy-based maps have been developed for all or parts of California (Joyner *et al.*, 1981; Tinsley and Fumal, 1985; VIC, 1993; Petersen *et al.*, 1997; Park and Elrick, 1998). More recently, Wills *et al.* (2000) presented the "first cut" of a geologic unit-based  $V_{S30}$  map for California that has found widespread usage in recent years. Further subdividing the original eight site classes used in the 2000 study, Wills and Clahan (2006) refined the urban areas (C. Wills, personal comm., 2006) of the earlier map by introducing 19 site classes based on a number of factors in addition to surficial geology—such as grain size, thickness of units, deep alluvium assumed for major basins, shallow depth for small and narrow basins, and degree of coarseness of sediments as a function of distance from mountain fronts.

While  $V_{S30}$  maps have traditionally been derived directly from geological maps, several studies (e.g., Romero and Rix, 2001; Wald and Allen, 2007; Yong *et al.*, 2008a; 2008b) have exploited newly available remote sensing data to develop site classification maps. As one of the earliest studies utilizing satellite imagery to account for site conditions, Romero and Rix (2001) used Landsat 7 Enhanced

Thematic Mapper Plus (ETM+) data to interpret the character of surficial sediments for estimating the potential of ground shaking in the Mississippi Embayment. Using Shuttle Radar Topography Mission digital elevation models (SRTM DEMs) (Farr *et al.*, 2007), Wald and Allen (2007) presented global  $V_{S30}$  maps based only on topographic slope. Most recently, Yong *et al.* (2008a; 2008b) performed semi-automated analyses of spectra (visible to thermal-infrared electro-magnetic frequency range) and 30-meter resolution DEM (stereoscopic correlation method) data produced from the Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER) satellite recordings to interpret remotely-sensed geologic and topographic information. Subsequently, these results were assembled into terrain types to estimate  $V_{S30}$  in a 60 km by 60 km area surrounding Islamabad, Pakistan. This study demonstrated that satellite data can help circumvent the problems associated with inconsistencies found in traditional geologic mapping. They also showed that the combination of multiple parameters, such as geology and topography, yields a more robust estimate than uniparametric models. However, results from earlier investigations (Yong *et al.*, 2005; 2006; 2007) indicated that a systematic and objective classification method that is both simple and yet robust for all terrain types is necessary for characterizing larger regions (e.g., Murray and Fonstad, 2007; Philips, 2007). To meet these requirements, we investigate the Iwahashi and Pike (2007) automated approach to develop a global terrain classification map that was also based on SRTM DEMs. As a basis for this study, we describe the details of our method where we explore the effectiveness of using a terrain-based model as the underlying framework for a  $V_{S30}$  proxybased approach.



Fig. 1. Terrain map of California (Iwahashi and Pike, 2007) with locations (black circles) of 853 V<sub>S30</sub> values.

## DATA

## Terrain-based Framework

Iwahashi and Pike (2007) (hereinafter as IP07) introduced an automated topography classification scheme to develop a globally consistent terrain classification map. The IP07 approach employed an algorithm that divides continuous topography (SRTM DEM) represented by elevations into a maximum possibility of 16 terrain classes through the use of a geometric signature consisting of three taxonomic criteria—slope gradient, local convexity and surface texture. The design of the topographic classification scheme is based on an unsupervised nested-means approach (Scripter, 1970) where the final output of terrain classes reflects the statistical properties of the input geometric variables rather than preset criteria. Unlike supervised algorithms, for which predetermined criteria for each output class do not adapt equally well to all locations and DEM resolutions, the unsupervised approach treats topography as a continuous random surface, independent of any spatial or morphological orderliness imposed by fluvial activity and other geomorphic processes. Since the resultant terrain classes are not determined a priori by training examples of target physiographic types, IP07 acknowledged that these classes are only approximately equivalent to such traditional terrain types as mountains, high hills, plains with hills, or tablelands in the manner of Hammond's (1964) taxonomy. Nevertheless, on the basis of subsequent map overlays and statistical analyses, IP07 showed that their U.S. terrain classification map closely resembles the classic physiographic divisions of Fenneman and Johnson (1946) and for more recent work, the geomorphic systems of Graf (1987). In our terrain framework, for example, the physiographic taxonomy established for the United States (Hammond, 1954; 1964) is used to relate the equivalent classes to the terrain category of California (Fig. 1). For the main part of this study, we use the same automated scheme and only adopt the California region out of the global dataset. In a later section, we discuss some results from preliminary tests, relating to use of an IP07 terrain map that covers the contiguous U.S.



Fig. 2. Barplot describing the combined frequency distribution of measured (black) and inferred (grey)  $V_{S30}$  values.

## V<sub>S30</sub> values

In developing our terrain-based  $V_{s30}$  proxy model, we use shear-wave velocity data provided from Walt Silva (last revised July 17, 2006; written comm., 2007) of Pacific Engineering and Analysis (PE&A). As part of the NGA (Next Generation Attenuation) Strongmotion Database (Chiou *et al.*, 2008), the  $V_{s30}$  data consists of both measured and inferred values, i.e., measurements from recorded waveforms and extrapolations from proxy- and/or map- based correlations, respectively. Hence, we will continue to purposefully refer to this mix of  $V_{S30}$  data as " $V_{S30}$  values" or "results" and not as "measurements" to avoid any misrepresentation about the nature of the source data. Of the 853 values in the database, 644 (75%) were inferred. Measured  $V_{S30}$  data were calculated from 201 (24%)  $V_S$ records with (minimum) 30 m depth profiles (Fig. 2). Other data used to infer  $V_{S30}$  include measurements are from six (0.7%)  $V_S$ profiles of less than 20 m and two (0.2%) borehole records extending to the 9 m depth. Furthermore, this sparse subset of measurement-derived  $V_{S30}$  values are incompatible due to the diversity of measurement methods used (Silva, personal comm., 2009). This limitation was previously noted by a number of authors (e.g., Wills, 1998; Boore, 2006; Moss, 2008). Biases between and imprecision associated with different methods control the overall epistemic uncertainties related to the results, and are directly attributable to the fundamental differences in principle that governs the respective geophysical method applied. For example, in a study comparing the variability of shear-wave velocity determinations from existing blind and comparative studies, Moss (2008) examined both the intramethod and intermethod variability and found a clear bias between invasive and noninvasive methods.

For  $V_{S30}$  values derived from proxy- and map-based inferences, concerns regarding data quality are generally related to the same issues as previously described for measured values. Despite these well-known issues, many of these data were used by Wills *et al.* (2000) to develop the Site Conditions Map of California and later by Wald and Allen (2007). In effect, the Silva (personal comm., 2007)  $V_{S30}$ database used in this study is an updated version of the  $V_{S30}$  values used by both Wills *et al.* (2000) and Wald and Allen (2007). Wills *et al.* (2000) used 556  $V_{S30}$  values; for their California map Wald and Allen (2007) used an updated database of 767 values as well as other values in active tectonic regions (e.g., Utah [204], Taiwan [387] and Italy [43]).

We use all 853  $V_{S30}$  results from the updated (2007) Silva  $V_{S30}$  database to develop our terrain-based  $V_{S30}$  proxy model. To maintain a framework consistent with the Wills *et al.* (2000) and Wald and Allen (2007) approaches, we also do not appraise the quality of the  $V_{S30}$  values used in our study other than acknowledging the lack of quality assurance in the  $V_{S30}$  data, and addressing the limitations of  $V_{S30}$  as a measure of site conditions later.

#### METHODS

#### Developing the Terrain-based V<sub>S30</sub> Proxy Model

After establishing our terrain-based framework, we compare available  $V_{S30}$  values with our terrain classification map. Using the spatial locations for the 853  $V_{S30}$  values, we determine the corresponding terrain-type class (Fig. 1). No  $V_{S30}$  values are associated with areas identified as our terrain category of Class 13, which fit the description of incised terraces (Fig. 2). Such a result might be expected because relatively few  $V_{S30}$  values are available at hard rock sites.  $V_{S30}$  values generally tend to be in major urban/metropolitan areas that are typically located on soft basin-type terrains (Fig. 2). Also, in the case of Class 13, the terrace terrain-type has the second-lowest ranking number (2,106 pixel cells) of occurrences in California representing <1% of the total cells (574,049) that make up the regionally extensive and morphologically diverse state (Yong *et al.*, 2010).

Following the association of  $V_{S30}$  values to the 15 site classes, we randomly select 556 values (same number used by Wills *et al.*, 2000) and calculate the mean  $V_{S30}$  values for each of the 15 terrain classes. Here, the portion of  $V_{S30}$  values assigned within each terrain class is based on the distribution of the original 853 values we find in each of the 15 terrain categories (Fig. 2). We use a standard cross-validation (non-replacement) method to compare our prediction to the remaining 297  $V_{S30}$  measurements by examining its mean-squared prediction error (MSPE):

$$\Delta = \frac{1}{N} \sum_{i=1}^{N} (V_{S30_i} - prediction_i)^2 \tag{1}$$

The routine, from random selection through cross-validation, is iterated one thousand times. At the end of these runs, we select the model with the lowest MSPE (14,414) as our best predictor of  $V_{S30}$  for each terrain type (Table 1).

		Freq.	Predicted
Terrain		in	<b>V</b> \$30
Туре	Description	CA (%)	(m/s)
1	Well dissected alpine summits, mountains, etc.	21	519
2	Large volcanoes, high block plateaus, etc.	2	393
3	Well dissected, low mountains, etc.	14	547
4	Volcanic fans, foot slopes of high block plateaus, etc.	19	459
5	Dissected plateaus, etc.	6	402
6	Basalt lava plains, glaciated plateaus, etc.	1	345
7	Moderately eroded plains of weak rocks, etc.	7	388
8	Desert alluvial slopes, volcanic fans, etc.	12	374
9	Well eroded plains of weak rocks, etc.	1	497
10	Valleys, till plains, etc.	<1	349
11	Eroded plains of weak rocks, etc.	2	328
12	Desert plains, delta plains, etc.	4	297
13	Incised terraces, etc.	<1	—
14	Eroded alluvial fans, till plains, etc.	<1	209
15	Dunes, incised terraces, etc.	1	363
16	Fluvial plains, alluvial fans, low-lying flat plains, etc.	8	246



### RESULTS

#### **Evaluation of Model Performances**

To assess our terrain-based  $V_{S30}$  proxy model beyond inferential statistics, we evaluate its performance compared to the geologic unitbased model (Wills *et al.*, 2000) and the topographic slope-based model (Wald and Allen, 2007). Since both of the latter approaches effectively employed the same  $V_{S30}$  data we use for the same area (state of California)—albeit, in varying quantities, i.e., Wills *et al.* (2000) and Wald and Allen (2007) relied on approximately 65% and 91% (respectively) of the values from the updated  $V_{S30}$ database—the commonality in data source reasonably supports the basis for our direct comparison of the different models. Using the "Spatial Join" function in the ESRI ArcInfo (version 9.3) application program, we associate the 853 sites to the seven categories of the Wills *et al.* (2000) model (data provided by C. Wills, written comm., 2006). On the basis of the site categories in the joined GISdatabase tables, we translate every instance of their categories from class letter designations to corresponding mean  $V_{S30}$  values (Wills *et al.*, 2000). For the Wald and Allen (2007) model, which is available from the USGS

(http://earthquake.usgs.gov/research/hazmaps/interactive/vs30/predefined.php; last accessed 18 July, 2008) as an evenly sampled xyz file, we develop a simple algorithm using the nearest neighbor approach (Clark and Evans, 1954) to associate their coordinate-and- $V_{S30}$ -value relations with the 853  $V_{S30}$  sites. To identify a straightforward and common parameter between the Wills *et al.* (2000) geologic unit-based categories and our own model, we use their reported mean  $V_{S30}$  values for each of their seven expanded NEHRP-based classes. For the Wald and Allen (2007) topographic slope-based categories, mean  $V_{S30}$  values were not reported (or readily available) so we use the assigned  $V_{S30}$  values in their downloadable text-file. We use the Warnes (2004) "bandplot" (locally-smoothed-mean and standard deviation) function and a customized subroutine—all in the open-source statistical environment "R" (last accessed 4 April, 2009; <u>http://www.r-project.org/</u>)—to consider the local mean and variance, and calculate the typical standard deviation ( $\sigma$  typical) of each model (Fig. 3). As a result, we find the Wills *et al.* (2000) model has the highest  $\sigma$  typical, approximately 195 m/s, while the Wald and Allen (2007) model and our model have approximately 142 m/s and 128 m/s, respectively. As part of our statistical analysis, we also calculate the median of the MSPE for each model and determine that all models perform significantly better than random.

As an extension of our efforts to develop a site characterization map for California, we explore the utility of our approach for a  $V_{S30}$  map of the contiguous U.S (Yong *et al.*, 2010). Like the California map (Fig. 1), the U.S.  $V_{S30}$  map (Fig. 4) is derived using the automated IP07 classification method that relies on the same taxonomic criteria (slope gradient, local convexity and surface texture) developed from geomorphometry to also identify 16 terrain types from the same 1-km spatial resolution (SRTM30) DEM. On the basis of the California model, we present our preliminary U.S. map by applying the same terrain classification approach and adopting the same  $V_{S30}$  relations determined for each terrain type in California. Although our database of 325  $V_{S30}$  observations (for source, see Data and Resources Section) measured outside of California is currently too sparse to permit a rigorous validation of the terrain-based



Fig. 3. Plots describing the smoothed mean and standard deviation for  $853 V_{s30}$  values based on both measured and inferred data (California) against the terrain-based (left), topographic slope-based (middle), and geological unit-based (right) estimation models. Red lines indicate the smoothed mean values, orange dashes indicate the one standard deviation from the mean and yellow dash indicates two standard deviations from the mean.



Fig. 4. Terrain map of the contiguous United States (Iwahashi and Pike, 2007) with locations (black circles) of  $325 V_{S30}$ measurement-based values, including 14 records from Canada. Areas of poor DEM quality were left unclassified and designated by the authors as black polygons.

model, we plot the locally-smoothed mean and standard deviation (Fig. 5) of the terrain-based and topographic slope models for a preliminary assessment of their performances. At the time of our investigation, no geology-unit based  $V_{S30}$  estimation model of the contiguous U.S. was readily available. When comparing the  $\sigma_{typical}$  of each model for their respective performances in California and the U.S. (Fig. 3 and 5), we note the changes ( $\Delta_{\sigma}$ ) in their  $\sigma_{typical}$  and find that the terrain-based model ( $\Delta_{\sigma} = +19$  m/s) is more stable than the topographic-slope based model ( $\Delta_{\sigma} = +65$  m/s). We also find an appreciable improvement in the predictive performance of the terrain-based model ( $\sigma_{typical} = 147$  m/s) over the topographic sloped-based model ( $\sigma_{typical} = 213$  m/s).



Fig. 5. Plots describing the smoothed mean and standard deviation for  $V_{S30}$  values based on 325 recorded measurements (contiguous U.S., outside of California) against the terrain-based (left) and topographic slope-based (right) estimation models. Red lines indicate the smoothed mean values, orange dashes indicate the one standard deviation from the mean and yellow dash indicates two standard deviations from the mean.

#### DISCUSSION

Our primary motives for developing the terrain-based model are effectively two-fold: to more fully exploit the nature of geomorphology as a proxy for site amplification, and to advance an approach that is completely objective with results firmly supported by robust statistical methods. We acknowledge that a significant portion (76%) of the data used to develop our model are based on inferred  $V_{S30}$  values and that such a strong dependence fundamentally contradicts the established principles of model development. However, it is necessary to depend on these proxy-based estimates because there are simply not enough openly-available measured values to meet standard statistical requirements. As previously indicated, the reliance on inferential data has precedent in state-of-art approaches, e.g., the geology-based and topographic slope-based models were also conditioned on effectively the same inference-based data. In the following section, we expand on these and other issues associated with our dependence on proxy-based models and the use of  $V_{S30}$  for estimating site amplification.

Although geomorphology is (in principle) an appropriate proxy for capturing the key characteristics of both geology and topography (Peterson, 1981; Harden, 1990; Ellis *et al.*, 1999; Saco *et al.*, 2006) and that morphometrics in terrain models meet the requirements for a consistent, objective, and quantitative framework that systematically encapsulates the parameters related to these near-surface properties (Oliver and Webster, 1986; Franklin, 1987; Dehn *et al.*, 2001; Minár and Evans, 2008; Pike *et al.*, 2009)—one initial concern is the role that terrain types play in our model. Because the proxies are by definition estimates of the predictors (i.e., terrain-type is the proxy for both geology and topography, which are in turn proxies for  $V_{S30}$ , which itself is a proxy for site amplification), this type of multi-tiered approach can inherently compound the uncertainties associated with each proxy. Despite these concerns, Iwahashi *et al.* (2010) examined results from multiple regression analyses of  $V_{S30}$  measurements at 1646 locations with topographic attributes (slope gradient, surface texture, and the logarithm of elevation) based on 50-m DEM and found that estimations of  $V_{S30}$  from a DEM are useful for earthquake vulnerability assessment.

One utility of the terrain-based approach is that material properties, which are not directly related to a single proxy, such as slope, can

be inferred. This is especially true when estimating  $V_{S30}$  in flat or steep topographic areas. For such instances, the combination of terrain nomenclature (Hammond 1954; 1964) and the IP07 classification scheme provides a basis for differentiating material properties that do not always directly associate with individual proxies, such as topographic slope. For example, in a volcanic plain with flat topographic slope and high shear-wave velocity, topographic slope (alone) would predict high amplification (low  $V_{S30}$ ). Furthermore, in active-tectonic regions, lower shear-wave velocity (high amplification) with steep topographic slope consisting of young sediments, a slope-based model would predict high  $V_{S30}$ . In both cases, geology plays the key role for distinguishing the exceptions in the relationship between the nature of topographic slope and  $V_{S30}$ . A manual analysis of the same regions (e.g., Modoc Plateau, Mendocino, etc.) in each model indicates significant differences in  $V_{S30}$  predicted (Fig. 1 and Wald and Allen, 2007). Actual observations of  $V_{S30}$  measurements from the regions are necessary to properly evaluate the performance of both models.

Clearly, the arbitrary nature of 30-meters as the key parameter for depth from the surface in the  $V_{S30}$  term is expected to have a strong first-order consequence on the effectiveness of the proxies. To address this issue, the use of more complicated parameterizations might be useful. Approaches worthy of further investigation, include the one-quarter wavelength method (Joyner *et al.*, 1981), or extra soil site considerations, e.g., sediment thickness coupled with shear-wave velocity (s<sub>L</sub>) first proposed by Seed *et al.* (1976) and recently reviewed by Lee and Trifunac (2010). Another possible approach is to circumvent  $V_{S30}$  as a proxy by determining if the IP07 terrain scheme can be used to directly estimate site amplification; this will require the use of records from both permanent (e.g., California Integrated Seismographic Network) and temporary (e.g., USArray) seismographic station sites in order to meet minimum spatial sampling requirements. For now, in the absence of a more robust site parameter, we follow accepted convention and proceed with  $V_{S30}$  as the proxy for material induced amplification.

Considering the uncertainties associated with the  $V_{S30}$  term, in addition to other vexing factors such as artifacts associated with spatially referenced data, not to mention circularity issues relating to the co-mingling of measured- and inferred- (map-based) values found in the  $V_{S30}$  database, we recognize that these limitations will tend to confound our dependence on strict adherence to our statistically-rigorous systematic approach. Rather, it is in the face of these uncertainties that we consider an objective and systematic approach based firmly on statistics to be most warranted when investigating the predictive capability of the IP07 terrain model to estimate  $V_{S30}$ .

By statistically determining that each model performs significantly better than random, we are able to confirm that all of the proxies have a strong correlation to  $V_{S30}$  (Yong et al., 2010). However, when we consider standard regression analysis methods to evaluate the performance of the models by themselves and against each other, we encounter a fundamental statistic-related problem inherent in all three models. The respective proxies of the geologic unit-, topographic slope- and terrain-based models do not readily fit in any formalized description of a statistical variable—i.e., although the proxies are defined as names (nominal characteristic), at the same time their associations to  $V_{S30}$  are strongly dependent on numeric values tied to predefined (discrete variable) ranges on a common continuous scale with meaningful hierarchical order (ordinal characteristic). Nevertheless, it is in the utility of the standard deviation  $(\sigma)$  where we find a meaningful, yet simple, statistical measure for evaluating the performance of each model and the comparisons against each other. Here, based on the 853 values from the Silva database, we look for the typical amount of dispersion ( $\sigma_{\text{typical}}$ ) in the  $V_{S30}$  values with respect to each model's proxy-based estimates. Because our terrain-based estimates depend on the average of the  $V_{S30}$ values for the respective terrain types, we choose the equivalent average  $V_{S30}$  values described by Will *et al.* (2000) for their eight respective classes to calculate the  $\sigma_{typical}$  of their geologic unit-based model. Although the topographic slope-based approach initially estimated  $V_{S30}$  on the basis of the median velocity of the subdivided NEHRP classifications, these results are not fully described in Wald and Allen (2007) and we find no means to calculate the average (nor median) of their  $V_{S30}$  estimates for each class because the complete data set is no longer available (Wald, written comm., 2009). In the face of these deficiencies, we instead rely on the latest and most-readily-available estimates from their USGS web-portal

(http://earthquake.usgs.gov/research/hazmaps/interactive/vs30/predefined.php; last accessed 18 July, 2008) for comparison to the Silva  $V_{S30}$  data. Because we consider their estimates to possess the most statistically advantageous traits on the basis of the continuous nature and sheer abundance (1,401) of the  $V_{S30}$  data used in their analyses, we expect Wald and Allen (2007) estimations to yield the least  $\sigma_{typical}$ , hence rank as the best performing model when compared against the discretized proxies in geologic unit- and terrain-based models. Instead, as noted earlier, we find our terrain-based model to yield the lowest  $\sigma_{typical}$  (approximately 128 m/s), follow closely by the topographic slope-based (approximately 142 m/s) and the geologic unit-based models (approximately 195 m/s), respectively.

Regardless of the complexity and/or innovation associated with automated and systematic approaches, the accuracy of the derived model is inherently limited by the quality of the input data. For the topography of California, IP07 used DEMs of spatial resolution at 30-arc-sec (approximately 1-kilometer grid spacing or pixel size) derived from recordings measured by the Shuttle Radar Topography Mission (SRTM) (Farr *et al.*, 2007). Although higher resolutions of SRTM DEMs were also available (e.g., the SRTM recordings were sampled over a grid at spatial resolution of 1-arc-second by 1-arc-second or approximately 30 m by 30 m), we instead choose to adhere to the proven IP07 technique which produced reasonably accurate classifications (Iwahashi and Pike, 2007). At the same time, the use of the same 30-arc-sec topography employed by Wald and Allen (2007) allows us to directly compare the two methods. By

comparison, the 1:250,000-scale maps used by Wills *et al.* (2000) have (in principle) the best-resolved details (pixel resolution of 125 m) with a detectable object-resolution of 250 m (Tobler, 1988).

#### CONCLUSION

To varying degrees of success, individual proxies such as mapped surface geology (Wills et al., 2000) and topographic slope (Wald and Allen, 2007) have been used to develop  $V_{S30}$  (the average shear-wave velocity in the upper 30 meters) predictions for California. This type of broad-brush (single-parameter) approach to site characterization has been demonstrated to be useful for some purposes, especially in the absence of measured  $V_{s30}$  values (Frankel et al., 2010; Thompson et al., 2010). However, clearly no single proxy can fully account for variations in material properties that control  $V_{S30}$ . In an attempt to improve site characterization, we develop a method based on geomorphometry (slope gradient, local convexity and surface texture) to better predict their material properties and thus estimate  $V_{s30}$ . We use results from the automated terrain classification method based on taxonomic criteria derived from geomorphometry to systematically identify 16 terrain types from a 1-km spatial resolution (SRTM30 data) DEM of California (Iwahashi and Pike, 2007). Using 853  $V_{S30}$  values, we assign  $V_{S30}$  values that are the average of values within each terrain type. We then compare these values to the single-proxy prediction models to determine relative performances. As expected, the overall results indicate that our approach performs better than the characterizations based on single proxies. In each case, we note that the reliance on inconsistent geologic data sets (Wills et al., 2000), or a simple parameter, such as the 1-km resolution SRTM30 DEM derived slope gradient (Wald and Allen, 2007), are less discriminating in accounting for variations in site conditions. Furthermore, our preliminary studies also indicate the terrain-based framework to be more effective than topographic-slope alone when comparisons are extended to the contiguous U.S. (Yong et al., 2010). Despite statistical tests that indicate our method to be, at worst case, a marginal improvement over the next-best method, we also note that by employing objective and systematic methods, as well as a multi-faceted framework that encapsulates geology and topographic-slope, our approach holds the most promise for improving the prediction of material (soil) properties when estimating site amplification. As new and spatially-distributed measurement-based  $V_{S30}$  data become available, we expect to be able to improve our predictive capabilities.

## DATA AND RESOURCES

GIS data of world-wide automated terrain classification map from SRTM 30 can be downloaded at <u>http://gisstar.gsi.go.jp/terrain/front\_page.htm</u>. The updated NGA Database Flat File (including measurement-based *V*<sub>530</sub> observations at 325 recording sites outside of California), referred in this article as the NGA (Next Generation Attenuation) Strong-motion Database, can be downloaded at <u>http://peer.berkeley.edu/nga/flatfile.html</u>.

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