Probabilistic Seismic Hazard Maps for Seattle: 3D Sedimentary Basin Effects, Nonlinear Site Response, and Uncertainties from Random Velocity Variations

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Outline of Talk

- Observations of 3D basin effects and site response for Seattle, WA region
- Modeling basin effects with 3D finite difference simulations; validating 3D velocity model
- Methodology for producing urban seismic hazard maps for Seattle using 3D simulations
- Evaluating effects of 3D random velocity variations on peak velocities and spectral accelerations from simulations









From 19 events M2.7-4.8

From M6.8 Nisqually EQ (dep=52 km)





1 Hz Amplification at stiff-soil sites in Seattle basin wrt rock site Amplification depends on direction to earthquake



3D finite difference modeling

- Used viscoelastic code of Pengcheng Liu (U.S. Bureau of Reclamation); 4th order in space, 2nd order in time
- Variable grid spacing with depth; finest spacing we used was 70m
- We used minimum Vs of 600m/s, similar to observed Vs30 of glacially-overridden soils
- Accurate to at least 1 Hz (9 grid points per wavelength)
- Validated 3D model with 5 earthquakes to date







600.0



Observed (black) and synthetic (red) velocity waveforms for The Nisqually earthquake (0.2-0.4 Hz)







Examples of observed and predicted basin surface waves (SW), 0.5-1.0 Hz

From Frankel et al. (2009)





Nisqually earthquake

M4.8 deep earthquake west of Seattle



science for a changing work

Putting the results of 3D groundmotion simulations into seismic hazard maps:

Seattle hazard maps; USGS Open-File Report 2007-1175 Frankel, Stephenson, Carver, Williams, Odum, Rhea



541 3D finite-difference simulations used in Seattle seismic hazard maps

- 458 simulations for earthquakes in Seattle fault zone (M6.6-M7.2)
- 9 simulations for earthquakes on Southern Whidbey Island fault
- 10 simulations for point sources on Cascadia subduction zone
- 48 simulations for shallow earthquakes: 8 azimuths, 3 distances and two depths (10 and 15 km)
- 16 simulations for deep earthquakes (50 km depth): 8 azimuths and 2 distances
- Calculated synthetics at 7236 sites, with 280m spacing
- Used about 7.8 million synthetic seismograms



Procedure to Make Urban Seismic Hazard Maps



PSHA= Probabilistic Seismic Hazard Assessment



Probabilistic seismic hazard with site and source dependent amplification and rupture directivity

Annual probability of having ground motion exceeding u_0 at site *i*:

$$P(u \ge u_0) \approx \sum_{M} \sum_{\text{source}_j} \text{rate}(M, \text{source}_j) P(u \ge u_0 | \text{site}_i, \text{source}_j, M)$$

For stiff-soil sites: $u=u_{rock}(M,D)A_{3D}(site_i,source_i)$

For soft-soil sites:

 $u = u_{rock}(M, D)A_{3D}(site_i, source_i)A_{soft}(site_i, PGA_{rock})$

Amp factor A_{3D} contains 3D basin effects and rupture directivity determined by 3D simulations for various scenarios

A_{soft} determined from Vs30 using Choi and Stewart (2005) empirical amplification factors



Float rupture zones along Seattle fault traces, do nine 3D simulations for each rupture zone (3 slip distributions, 3 hypocenters)



Two scenarios for Seattle fault earthquakes M6.6



Used kinematic description of rupture on fault surface





 Epicenters used in simulations to determine azimuthal dependence of amplification for hazard from background earthquakes

Black shallow: 10, 15 km depth Red: 50 km depth

Distances chosen from dominant values from hazard deaggregation

Possible configurations for rupture zone of great Cascadia Earthquake



Figure from Petersen et al. (2002)





Point sources used to quantify amplification expected from great Cascadia earthquakes



3.5

2.5 1.5

20000

5000

00

E-W Distance (m)

5000

00

E-W Distance (m)

20000

1 Hz amplification Cascadia point sources, from **3D** simulations

Aleatory Uncertainty

To calculate hazard from each source grid cell or each rupture scenario we applied aleatory sigma from empirical GMPE's for firm-rock site condition

We also tried applying aleatory sigma to median of 1 sec S.A. for the 9 scenarios for each rupture zone. This produces very similar results at 10% and 2% PE as first approach.

Our approach ensures hazard values of new maps will be consistent with NSHM's for firm-rock sites outside of Seattle basin



1 Hz Spectral Acceleration (%g) with 2% chance of being exceeded in 50 years



one of 2002 national seismic hazard maps; Firm-rock site condition Using 3D simulations with basin effects and directivity Using 3D simulations and nonlinear ampl. for fill/alluvium





Map of thickness of fill/alluvium, determined by Susan Rhea using compilation of borehole data by Kathy Troost (Univ. of WA)

This map of thickness was then used to make a map of Vs30, using an average Vs profile for fill/alluvium sites

Amplification at soft-soil sites determined from Vs30 using Choi and Stewart (2005) empirical amplification factors





1 Hz S.A. with 50% Probability of Exceedance in 50 Years



We need to include basin amplification terms in building codes 1 Hz S.A. (%g) with 2% Prob. Of Exceedance in 50 Years





From 2002 national seismic New map with basin effects, hazard maps and NEHRP amplification rupture directivity, factors based on Vs30 and nonlinear soil response from surficial geology at soft-soil sites



What are the effects of realistic 3D random spatial variations of Vs and Vp on the ground motions in the 3D simulations?



Colors represent shear wave velocity variations



- Von Karman correlation function, stddev of Vs = 10% in top 1.3 km; 5% from 1.3-10.8 km depth; correlation distance of 5 km
- Randomness in Vs is fractal for length scales less than about 30 km (equal variance of Vs over equal log increments of wavelength
- Vs and Vp variations are correlated
- Minimum Vs = 500 m/s; minimum of mean Vs = 600 m/s
- This stddev for shallow basin Vs is consistent with variations found in borehole studies (e.g., ROSRINE; Thelan et al., 2006)





Figure from Thelan et al. (2006) showing variability in Vs30 in Los Angeles basin and San Gabriel Valley as a function of site separation; from borehole measurements compiled by Wills and Silva (1998) and Gibbs et al., (2000,2001) Power spectrum P(k) of random variations for von Karman correlation function (with order m=0)

 $P(k) = C (1 + (ka)^2)^{-3/2}$

k is radial wave number for 3D medium

a is correlation distance

C is a constant





Shear-wave velocity (m/s) at 1.4 km depth

With 3D random variations Original model 60000 60000 50000 50000 3200.00 **NS Distance (m)** 30000 20000 20000 NS Distance (m) 2922.22 40000 2644.44 2366.67 2088.89 1811.11 30000 1533.33 1255.56 977.78 700.00 20000 20000 10000 10000 0 0 20000 40000 60000 20000 60000 0 40000 0 EW Distance (m) EW Distance (m)

Used 10% std dev in top 1.3 km, 5% std dev from 1.3 to 10.8 km depth, von Karman correlation function (wide range of scale lengths); Hurst exponent = 0



PGV's (m/s) for simulations of M6.7 earthquake on Seattle fault



PGV's plotted are geometrical mean of PGV of two horizontal components



Ratio of PGV's between randomized and original models



Random variations in Vs produce epistemic uncertainty in ground motions



Random variations in Vs strongly affect amplitudes of basin surface waves in simulations

All synthetic seismograms are NS velocity






Observations of variability of basin surface waves (SW) across a 500m aperture array in San Leandro, CA (transverse acceleration records from M4.1 Alamo earthquake, filtered between 0.5 and 1.0 Hz)

Perhaps this is caused by random spatial variations of Vs in East Bay sediments



PGV in two East-West lines across basin





 Random variations in seismic velocity tend to reduce PGV and spectral accelerations in the direction of maximum forward directivity



Ratios of PGV's between randomized and original models Different seeds for random variations; same slip distribution on fault



Dashed ellipse is approximate location of Seattle basin

Rectangles are areas of reduced average PGV in forward rupture direction caused by scattering from random variations



Ratio of 1.0 sec Spectral Accelerations between randomized and original models

Ratio of 3.0 sec Spectral Accelerations between randomized and original models





Ratios of spectral accelerations are taken from geometrical mean of spectral acceleration at each site over two horizontal components

Note low ratios (deamplification) north of ends of fault at sides of basin Higher ratios in center of basin and outside the basin



Simulations using eastern segment of fault: Shifting hypocenter changes location of deamplification





Histograms of ratios of 1.0 sec S.A. between random and original models

1.0 sec Spectral Acceleration



Excess of low values caused by decrease of amplitude at updip sites



- Random Vs variations produce epistemic uncertainty in spectral acceleration that is a significant portion of the so-called "aleatory" uncertainty of the misfit of GMPE's to data
- Standard deviation (In units) for stiff-soil sites:

	1.0 sec S.A.	3.0 sec S.A
From random variations of Vs	0.34	0.27
Campbell and Bozorgnia (2008)	0.62	0.65



Random variations produce larger stddev (sigma) at basin sites than hanging wall sites because basin surface waves are more sensitive to random variations than are steeply propagating S-waves in hanging wall.

May be tendency to reduce median amplitude of hanging wall sites





Why Should We Care?

- Earthquake scenarios in 3D models without random variations in Vs may overestimate areas with focused basin surface waves; may overestimate PGV and 1.0 and 3.0 sec S.A. in forward rupture direction for sites and underestimate amplitudes in other directions
- Probabilistic hazard maps such as Seattle maps using hundreds of scenarios mitigate this problem
- Random variations in Vs (stddev of 10%) can produce localized amplification of a factor of two in PGV and 1.0 and 3.0 sec S.A. over distances of a km or so; could explain some cases of localized differences in damage from earthquakes



Why Should We Care?

- For PSHA: random variations in Vs produce significant epistemic (modeling) uncertainty of ground-motion values that will affect calculations of hazard; epistemic In sigma of 0.3 for basin sites (for PGV, 1.0 and 3.0 sec S.A.), a substantial portion of observed sigma from GMPE misfit of data
- We need to better assess the variability of basin surface waves caused by small-scale fluctuations of Vs, using array observations and simulations, to improve our estimates of ground-motion uncertainty (sigma) for urban seismic hazard maps and to provide synthetic seismograms that capture the variability of basin surface waves for the design of long-period buildings

